

# Large eddy simulations of stratified turbulence: the dynamic Smagorinsky model

Sina Khani<sup>1,†</sup> and Michael L. Waite<sup>1</sup>

<sup>1</sup>Department of Applied Mathematics, University of Waterloo, Waterloo, Ontario, N2L 3G1, Canada

(Received 29 September 2014; revised 25 February 2015; accepted 27 April 2015)

The dynamic Smagorinsky model for large eddy simulation (LES) of stratified turbulence is studied in this paper. A maximum grid spacing criterion of  $\Delta/L_b < 0.24$  is found in order to capture several of the key characteristics of stratified turbulence, where  $\Delta$  is the filter scale and  $L_b$  is the buoyancy scale. These results show that the dynamic Smagorinsky model needs a grid spacing approximately twice as large as the regular Smagorinsky model to reproduce similar results. This improvement on the regular Smagorinsky eddy viscosity approach increases the accuracy of results at small resolved scales while decreasing the computational costs because it allows larger  $\Delta$ . In addition, the eddy dissipation spectra in LES of stratified turbulence present anisotropic features, taking energy out of large horizontal but small vertical scales. This trend is not seen in the non-stratified cases, where the subgrid-scale energy transfer is isotropic. Statistics of the dynamic Smagorinsky coefficient  $c_s$  are investigated; its distribution is peaked around zero, and its standard deviations decrease slightly with increasing stratification. In line with previous findings for unstratified turbulence, regions of increased shear favour smaller  $c_s$  values; in stratified turbulence, the spatial distribution of the shear, and hence  $c_s$ , is dominated by a layerwise pancake structure. These results show that the dynamic Smagorinsky model presents a promising approach for LES when isotropic buoyancy-scale resolving grids are employed.

**Key words:** stratified turbulence, turbulence modelling, turbulence simulation

## 1. Introduction

An alternative approach to direct numerical simulation (DNS) is large-eddy simulation (LES), in which scales larger than the grid spacing  $\Delta$  are resolved, but subgrid-scale (SGS) effects are parametrized. A common and practical SGS scheme is the Smagorinsky (1963) model, in which the deviatoric part of the SGS momentum tensor  $\tau^r$  is expressed in terms of the filtered rate of strain  $\bar{s}_{ij} = 1/2(\partial \bar{u}_i/\partial x_j + \partial \bar{u}_j/\partial x_i)$ , as follows (using the notation of Pope 2000):

$$\tau_{ij}^r(\mathbf{x}, t) = -2\nu_r(\mathbf{x}, t)\bar{s}_{ij}(\mathbf{x}, t), \quad (1.1)$$

<sup>†</sup> Email address for correspondence: [sinakhani@uwaterloo.ca](mailto:sinakhani@uwaterloo.ca)

where  $\bar{\mathbf{u}} = (\bar{u}, \bar{v}, \bar{w})$  is the filtered velocity field. The eddy viscosity coefficient  $\nu_r(\mathbf{x}, t)$  is defined by the following model:

$$\nu_r = c_s \Delta^2 \bar{S}, \quad (1.2)$$

where  $c_s$  is the Smagorinsky coefficient, and  $\bar{S} = (2\bar{s}_{ij}\bar{s}_{ij})^{1/2}$ . (Here we define  $c_s$  without the squared power, as in Germano 1992, Ghosal *et al.* 1995 and Pope 2004, for example.) A constant value of  $c_s \approx (0.17)^2$ , which was suggested by Lilly (1967), did not work particularly well in complex turbulent flows; as a result, Siegel & Domaradzki (1994) investigated various ranges of  $c_s$  from  $(0.13)^2$  to  $(0.24)^2$  for different turbulent flows. Indeed, there is no clear approach for selecting  $c_s$  in complex turbulent flows. The dynamic Smagorinsky model is a method proposed by Germano *et al.* (1991), in which a temporally and spatially varying  $c_s$  is computed by applying a second filter  $\tilde{\Delta}$  and assuming a self-similar inertial subrange between two filter scales. An improvement by Lilly (1992) yielded a method to find  $c_s$  using the resolved fields.

In the last few years, due to high cost of DNS (e.g. Almalkie & de Bruyn Kops 2012; Bartello & Tobias 2013), there has been increased interest in using LES for computational studies of stratified turbulence (e.g. Remmler & Hickel 2012; Paoli *et al.* 2013; Khani & Waite 2014). In stratified turbulence, we also need to model the SGS density flux  $\mathbf{h}$ , which is related to the filtered perturbation density  $\bar{\rho}(\mathbf{x}, t)$  via

$$h_j(\mathbf{x}, t) = -\frac{2}{Pr_t} \nu_r(\mathbf{x}, t) \frac{\partial \bar{\rho}(\mathbf{x}, t)}{\partial x_j}, \quad (1.3)$$

where  $Pr_t$  is the turbulent Prandtl number. Khani & Waite (2014) have investigated the performance of two classical LES approaches, the Smagorinsky (1963) and Kraichnan (1976) models, in LES of stratified turbulence where the filter width  $\Delta$  is larger than the Ozmidov scale. This study found a necessary criterion on  $\Delta$  for LES to capture the fundamental dynamics of stratified turbulence, including a cascade to small scales (Lindborg 2006), a horizontal wavenumber energy spectrum with a slope around  $-5/3$  (e.g. Lindborg 2006; Brethouwer *et al.* 2007; Almalkie & de Bruyn Kops 2012; Bartello & Tobias 2013), and a breakdown of the layerwise structure into Kelvin–Helmholtz (KH) instabilities (Laval, McWilliams & Dubrulle 2003; Waite 2011; Bartello & Tobias 2013; Khani & Waite 2014). These criteria are related to the buoyancy scale

$$L_b = 2\pi \frac{u_{rms}}{N}, \quad (1.4)$$

where  $u_{rms}$  is the root-mean-square velocity and  $N$  is the buoyancy frequency. The buoyancy scale is known to be the thickness of the layers that emerge in stratified turbulence when viscous effects are not dominant (Waite & Bartello 2004; Brethouwer *et al.* 2007). As concluded by Khani & Waite (2014), the performance of the Smagorinsky model in LES of stratified turbulence is disappointing since it requires three times the resolution compared to the Kraichnan model in order to adequately capture some of the key features of stratified turbulence ( $\Delta/L_b < 0.17$  for the Smagorinsky LES versus  $\Delta/L_b < 0.47$  for the Kraichnan LES). On the other hand, the applicability of the Kraichnan LES is limited to problems with idealized periodic boundary conditions, so it is not appropriate for turbulent flows near boundaries or in complex geometries; note, however, that physical space eddy viscosity models also face challenges near boundaries (e.g. Meneveau & Katz 2000; Pope 2000; Wan & Porté-Agel 2011; Lu & Porté-Agel 2014).

Despite its popularity for geophysical flows, problems with the Smagorinsky model, including excessive dissipation near the filter scale, are widely known (Germano *et al.* 1991; Ghosal *et al.* 1995; Meneveau & Katz 2000; Pope 2000). In other fields such as engineering flows, wall-bounded turbulence, and boundary layers, the dynamic Smagorinsky model is much more widely used (e.g. Jiménez & Moser 2000; Meneveau & Katz 2000; Pope 2000; Wan & Porté-Agel 2011; Lu & Porté-Agel 2014; Smith & Porté-Agel 2014), but its applicability to stratified turbulence has not been investigated. In this paper, we perform and analyse LES of stratified turbulence using the dynamic Smagorinsky model, in which the main goal is to decrease the computational costs of the Smagorinsky model in resolving the dynamics of stratified turbulence. We investigate the maximum filter scale  $\Delta$ , which allows for capturing some of the fundamental features of stratified turbulence, including a cascade to small scales, a  $-5/3$  horizontal wavenumber energy spectrum and breakdown of layers into KH instabilities. Also, we investigate the anisotropy of the SGS energy transfer, the statistics of  $c_s$  at different buoyancy frequencies and numerical resolutions, and the relationship between  $c_s$  and the resolved dynamics.

In § 2, background on stratified turbulence and the dynamic Smagorinsky model are reviewed. The numerical approach and methodology are described in § 3. Section 4 presents results, and conclusions are given in § 5.

## 2. Background

The non-dimensional filtered Navier–Stokes equations under the Boussinesq approximation are

$$\frac{\partial \bar{u}_i}{\partial t} + \frac{\partial}{\partial x_j} (\bar{u}_i \bar{u}_j) = -\frac{\partial \bar{p}}{\partial x_i} - \frac{1}{Fr_\ell^2} \bar{\rho} \mathbf{e}_z - \frac{\partial \tau_{ij}^r}{\partial x_j} + \bar{f}_i, \quad (2.1)$$

$$\frac{\partial \bar{u}_j}{\partial x_j} = 0, \quad (2.2)$$

$$\frac{\partial \bar{\rho}}{\partial t} + \frac{\partial}{\partial x_j} (\bar{\rho} \bar{u}_j) - \bar{w} = -\frac{\partial h_j}{\partial x_j}, \quad (2.3)$$

where  $\bar{\mathbf{u}}$ ,  $\boldsymbol{\tau}$ ,  $\mathbf{h}$ ,  $p$  and  $\mathbf{f}$  are the filtered velocity, deviatoric SGS momentum flux, SGS density flux, perturbation pressure and velocity forcing fields, respectively, and the Froude number  $Fr_\ell = u/N\ell$  is defined based on a velocity scale  $u$  and a length scale  $\ell$ . We neglect the molecular viscosity and diffusion because of the assumption of large Reynolds number.

The dynamic Smagorinsky model applies a second test filter  $\tilde{\Delta} > \Delta$  to the momentum equation (2.1). It is common to choose  $\tilde{\Delta} = 2\Delta$  (e.g. Meneveau & Katz 2000; Pope 2000). The test-filtered momentum equation is given by (e.g. Pope 2000)

$$\frac{\partial \tilde{\tilde{u}}_i}{\partial t} + \frac{\partial}{\partial x_j} (\tilde{\tilde{u}}_i \tilde{\tilde{u}}_j) = -\frac{\partial \tilde{\tilde{p}}}{\partial x_i} - \frac{1}{Fr_\ell^2} \tilde{\tilde{\rho}} \mathbf{e}_z - \frac{\partial T_{ij}^r}{\partial x_j} + \tilde{\tilde{f}}_i, \quad (2.4)$$

where  $T_{ij}^r = T_{ij} - 1/3 T_{rr} \delta_{ij}$ , and

$$T_{ij} = \widetilde{\bar{u}_i \bar{u}_j} - \tilde{\tilde{u}}_i \tilde{\tilde{u}}_j \quad (2.5)$$

is the sub-test-filter-scale momentum tensor, which needs to be modelled due to the unknown term  $\widetilde{\bar{u}_i \bar{u}_j}$ . Similar to  $\boldsymbol{\tau}^r$  in (1.1), a closure model for  $\mathbf{T}^r$  using the

Smagorinsky approach could be defined as

$$T_{ij}^r = -2c_s \tilde{\Delta}^2 \tilde{\tilde{S}}_{ij}, \quad (2.6)$$

where the Smagorinsky coefficient  $c_s$ , which is now allowed to depend on position and time, should be the same as that in (1.2) because of self-similarity and the scale-independent assumption (see e.g. Porté-Agel, Meneveau & Parlange (2000) for the scale-dependent dynamic SGS model). Applying the test filter  $\tilde{\Delta}$  to the SGS momentum tensor  $\tau_{ij}$  and its Smagorinsky model in (1.1), and then subtracting them from the test SGS momentum tensor  $T_{ij}$  and its Smagorinsky model (2.6), respectively, yield (e.g. Meneveau & Katz 2000; Pope 2000)

$$L_{ij} = T_{ij} - \tilde{\tau}_{ij} = \widetilde{\tilde{u}_i \tilde{u}_j} - \tilde{\tilde{u}}_i \tilde{\tilde{u}}_j, \quad (2.7)$$

$$M_{ij} = \Delta^2 \tilde{\tilde{S}}_{ij} - \tilde{\Delta}^2 \tilde{\tilde{S}}_{ij}, \quad (2.8)$$

which are the resolved stress tensor and the Germano rate of strain tensor, respectively, and are related by the equation

$$L_{ij}^r = 2c_s M_{ij}. \quad (2.9)$$

Since  $L_{ij}^r$  and  $M_{ij}$  are known from (2.7) and (2.8), the only unknown in (2.9) is  $c_s$  and as a result, the system of equation is extremely overdetermined. Lilly (1992) suggests a least-squares approach to get  $c_s$ , which gives

$$c_s = \frac{1}{2} \frac{L_{ij}^r M_{ij}}{M_{ij} M_{ij}}. \quad (2.10)$$

It is worthwhile noting the physical interpretation of (2.10), which is that  $c_s$  is characterized by the amount of dissipation that is generated by projection of the resolved stress  $L_{ij}^r$  on the Germano rate of strain  $M_{ij}$  (e.g. Jiménez & Moser 2000; Meneveau 2012).

When  $L_{ij}^r$  and  $M_{ij}$  are not coaxial,  $c_s$  is negative, which might be interpreted as backscatter (e.g. Germano *et al.* 1991). In this situation, any attempts to model the stress tensor as proportional to the rate of the strain tensor will fail (e.g. Jiménez & Moser 2000). Indeed, negative eddy viscosity and diffusivity coefficients in (2.1) and (2.3), respectively, inevitably lead to numerical instabilities (e.g. Domaradzki, Liu & Brachet 1993; Ghosal *et al.* 1995; Meneveau & Katz 2000; Pope 2000; Remmler & Hickel 2012). There are two common approaches to avoiding negative  $c_s$ : clipping  $c_s$  by setting negative values to a non-negative threshold (e.g. 0) and averaging  $c_s$  over homogeneous directions (e.g. Lilly 1992; Ghosal *et al.* 1995; Lesieur & Métais 1996; Piomelli 1999; Meneveau & Katz 2000; Pope 2000; Porté-Agel *et al.* 2000; Lu & Porté-Agel 2014). In this study we apply the first approach: negative  $c_s$  values are set to zero, but otherwise  $c_s^+$  is fully time- and space-dependent, i.e.

$$c_s^+(\mathbf{x}, t) = \begin{cases} c_s(\mathbf{x}, t) & c_s(\mathbf{x}, t) \geq 0, \\ 0 & c_s(\mathbf{x}, t) < 0. \end{cases} \quad (2.11)$$

For comparison, we will also investigate the performance of the volume-averaged  $c_s$  in our simulations.

In this work we consider strongly stratified turbulence, i.e. with  $Fr_\ell \ll 1$ , in which the large-scale dynamics are dominated by quasi-horizontal vortical motions rather than gravity waves (Riley & de Bruyn Kops 2003; Waite & Bartello 2004; Lindborg 2006). Such flows are known to have a layered structure with layer thickness around  $L_b$  (Waite & Bartello 2004) and a direct cascade of energy to small scales (Lindborg 2006). Simulations with isotropic resolution of  $L_b$  show that the layers break up into KH instabilities on the buoyancy scale (Laval *et al.* 2003; Brethouwer *et al.* 2007; Waite 2011; Bartello & Tobias 2013; Waite 2014); as a result, the horizontal wavenumber kinetic energy spectrum has an approximately  $-5/3$  power law (Lindborg 2006; Brethouwer *et al.* 2007; Almalkie & de Bruyn Kops 2012; Bartello & Tobias 2013) with possibly a bump at  $L_b$  (Laval *et al.* 2003; Brethouwer *et al.* 2007; Waite 2011). These characteristics of stratified turbulence – a downscale cascade,  $-5/3$  horizontal energy spectrum and KH instabilities at the buoyancy scale – only occur when the buoyancy Reynolds number  $Re_b = \epsilon/\nu N^2 \gg 1$  and the horizontal Froude number  $Fr_h \ll 1$ , where  $Fr_h = u_{rms}/Nl_h$  (Riley & de Bruyn Kops 2003; Brethouwer *et al.* 2007; Almalkie & de Bruyn Kops 2012; Bartello & Tobias 2013). Note that when  $l_h$  is defined using the Taylor hypothesis, it follows that  $Re_b = Fr_h^2 Re$ , where  $Re = u_{rms}l_h/\nu$  (as in Brethouwer *et al.* 2007 and Bartello & Tobias 2013, for example). When the dissipation is too strong for a given stratification, e.g. when  $Re_b < 1$ , the energy cascade to small horizontal scales is eliminated, the horizontal wavenumber energy spectrum becomes very steep, and KH instabilities are suppressed (e.g. Brethouwer *et al.* 2007; Bartello & Tobias 2013).

### 3. Methodology

Forced stratified turbulence is studied in this paper, where the forcing term is applied to the rotational part of the horizontal velocity field, i.e. vortically forced stratified turbulence; see Herring & Métais (1989), Waite & Bartello (2004), Waite (2011), Khani & Waite (2014) for more details. Idealized simulations in a cubic box of side  $L = 2\pi$  are considered. The sharp spectral filter is employed, where the relationship between the cutoff wavenumber  $k_c$  and the grid spacing  $\Delta$  is given by

$$\Delta = \frac{\pi}{k_c}. \quad (3.1)$$

Spatial derivatives are discretized using the spectral transform method, where the two-thirds rule (Orszag 1971) is applied in each direction for the elimination of aliasing errors. Hence, the cutoff wavenumber  $k_c$  is related to the resolution as

$$k_c = \pi \frac{2n}{3L}, \quad (3.2)$$

(i.e.  $2/3$  of the Nyquist wavenumber), where  $n$  is the number of grid points in the  $x$ ,  $y$  and  $z$  directions. We can use (3.2) in (3.1) to get the effective grid spacing  $\Delta = 3L/2n$ , which is used in the eddy viscosity equations. For time advancement, the explicit third-order Adams–Bashforth scheme is employed.

Simulations are initialized with low-level random noise and spun up at low resolution ( $n = 256$ ) and with hyperviscosity to  $t = 300$ . Simulations are then continued at low ( $n = 256$ ) and high ( $n = 512$ ) resolution with the dynamic Smagorinsky subgrid-scale model; a similar approach is considered in Waite & Bartello (2004), Waite (2011), Khani & Waite (2014), for example. Additional simulations with the regular Smagorinsky model (described in Khani & Waite 2014) are also considered

	$N$	$n$	$k_b$	$k_c$	$k_o$	$\langle\epsilon\rangle$	$\langle E(t)\rangle$	$Fr_h$	$k_c/k_b$	$\Delta/L_b$	$\Delta t$
Dynamic Smagorinsky											
d5N0	0	512		168		$1.5 \times 10^{-4}$	0.0058	$\infty$			0.006
d5N2	2	512	31	168	271	$1.1 \times 10^{-4}$	0.0043	0.013	5.42	0.092	0.009
d5N4	4	512	57	168	756	$1.1 \times 10^{-4}$	0.0049	0.006	2.95	0.169	0.008
d5N6	6	512	80	168	1352	$1.2 \times 10^{-4}$	0.0056	0.003	2.10	0.239	0.007
d2N0	0	256		83		$1.4 \times 10^{-4}$	0.0053	$\infty$			0.018
d2N2	2	256	30	83	273	$1.1 \times 10^{-4}$	0.0044	0.012	2.77	0.180	0.018
d2N4	4	256	54	83	777	$1.1 \times 10^{-4}$	0.0055	0.005	1.53	0.327	0.016
d2N6	6	256	72	83	1327	$1.2 \times 10^{-4}$	0.0070	0.003	1.15	0.435	0.015
d2N12	12	256	127	83	3679	$1.3 \times 10^{-4}$	0.0090	0.001	0.65	0.763	0.010
Smagorinsky											
S7N2	2	768	32	254	286	$9.8 \times 10^{-5}$	0.0039	0.012	7.94	0.063	0.006
S5N2	2	512	31	168	262	$1.2 \times 10^{-4}$	0.0042	0.014	5.42	0.092	0.009
S2N2	2	256	32	83	282	$1.0 \times 10^{-4}$	0.0040	0.013	2.59	0.193	0.018

TABLE 1. List of numerical simulations.

for comparison. Four different buoyancy frequencies of  $N = 2, 4, 6$  and  $12$ , along with an unstratified case ( $N = 0$ ), are considered. Froude numbers vary from  $0.001$  to  $\infty$  to cover both strongly stratified (e.g. Riley & de Bruyn Kops 2003; Waite & Bartello 2004; Hebert & de Bruyn Kops 2006a; Lindborg 2006; Brethouwer *et al.* 2007; Waite 2011; Khani & Waite 2013, 2014) and unstratified turbulence. The typical kinetic energy dissipation rate  $\epsilon$  is around  $10^{-4}$ , which gives a forcing time scale  $t_f \sim 10$ , when combined with the forcing wavenumber  $k_f = 3$ . The turbulent Prandtl number  $Pr_t = 1$  and  $u_{rms} = \sqrt{\langle E(t) \rangle}$  (e.g. Khani & Waite 2014), where the angle brackets  $\langle \cdot \rangle$  denote time-averaging over  $375 \leq t \leq 450$ . Table 1 shows parameters and averaged quantities for the dynamic Smagorinsky LES in this paper, where  $k_o = (N^3/\epsilon)^{1/2}$  is based on  $\langle \epsilon \rangle$  and  $k_b$  is based on  $u_{rms}$ .

As in Khani & Waite (2014), we consider the ability of the LES model to capture fundamental features of stratified turbulence at different resolutions. We have identified three features of stratified turbulence that are known to be suppressed in simulations when dissipation is too strong for a given stratification: an approximately  $-5/3$  energy spectrum in the horizontal wavenumber; KH instabilities and billows on the stratified turbulence layers, with associated regions of small and negative Richardson number; and shallowing, or bumps, in the energy spectrum at the buoyancy scale, associated with the injection of energy from the KH instabilities. These characteristics have been found in many numerical studies of stratified turbulence, including DNS and hyperviscosity experiments: Lindborg (2006), Brethouwer *et al.* (2007), Almalkie & de Bruyn Kops (2012) and Bartello & Tobias (2013) have found approximately  $-5/3$  energy spectra; Laval *et al.* (2003), Brethouwer *et al.* (2007) and Waite (2011, 2014) have discussed KH instabilities and associated bumps in the energy spectra. The recent high-resolution DNS by Bartello & Tobias (2013) also found KH instabilities and stratification-dependent bumps in the compensated energy spectra. These are not the only properties of stratified turbulence, but they are important, and they are clearly suppressed when dissipation is too strong; as a result, we have chosen to focus on these phenomena to identify a resolution criterion for LES. For LES with the Smagorinsky (1963) or Kraichnan (1976) models, it seems only necessary that  $L_b$  be sufficiently resolved (Khani & Waite 2014).

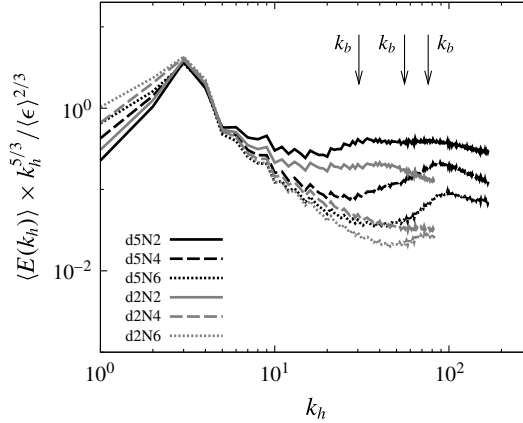


FIGURE 1. The averaged compensated horizontal energy spectra. The spectra are averaged over  $375 \leq t \leq 450$ . From left to right, arrows show buoyancy wavenumbers  $k_b$  that correspond to buoyancy frequencies  $N=2, 4$ , and  $6$ , respectively.

## 4. Results and discussion

### 4.1. Buoyancy scale effects on the dynamic Smagorinsky model

The compensated horizontal wavenumber energy spectra, in which the horizontal spectra are normalized by  $k_h^{-5/3} \langle \epsilon \rangle^{2/3}$ , are shown in figure 1. The advantage of using the compensated spectra is that constant horizontal spectra at intermediate wavenumbers imply a spectral slope of  $-5/3$ , in line with the stratified turbulence inertial subrange theory of Lindborg (2006). The high-resolution case with  $N=2$  shows an almost constant normalized spectrum over  $6 \lesssim k_h \lesssim 30$  along with a wide bump around  $k_b \sim 30$ . The lower-resolution simulation with the same stratification exhibits a steeper spectrum, in which the bump around the buoyancy scale is weakly resolved. Increased stratification steepens the compensated horizontal wavenumber energy spectra, to the extent that the high-resolution cases with  $N=4$  and  $6$  do not clearly show constant inertial subranges but resolve (small) bumps around  $k_b \sim 60$  and  $80$ , respectively. In addition, the low-resolution LES with larger stratification do not seem to capture any inertial subrange or bumps. In agreement with LES of stratified turbulence in Khani & Waite (2014), these results suggest that resolving a clear constant inertial subrange in the compensated horizontal wavenumber spectrum depends on the resolution and buoyancy frequency. As a result, stronger stratification may need higher resolution in LES to capture an inertial subrange.

Khani & Waite (2014) have shown that in the Kraichnan and Smagorinsky LES, capturing dynamics of stratified turbulence depends on the ratio  $\Delta/L_b$ . Only for sufficiently small values of  $\Delta/L_b$  – below  $0.47$  for Kraichnan and  $0.17$  for Smagorinsky – were the small-scale features of stratified turbulence, including KH instabilities and locally small values of the Richardson number, captured. To evaluate the relevance of this criterion for the dynamic Smagorinsky model, we consider the horizontal vorticity fields as well as distributions of the local Richardson number in our stratified simulations.

Figure 2 shows the  $y$  component of vorticity  $\bar{\omega}_y$  in the  $x$ – $z$  plane for the high-resolution simulations with different buoyancy frequencies at  $y = 0.25$  and  $t = 450$ . Unlike the unstratified case (figure 2a), the stratified cases are layered in the vertical



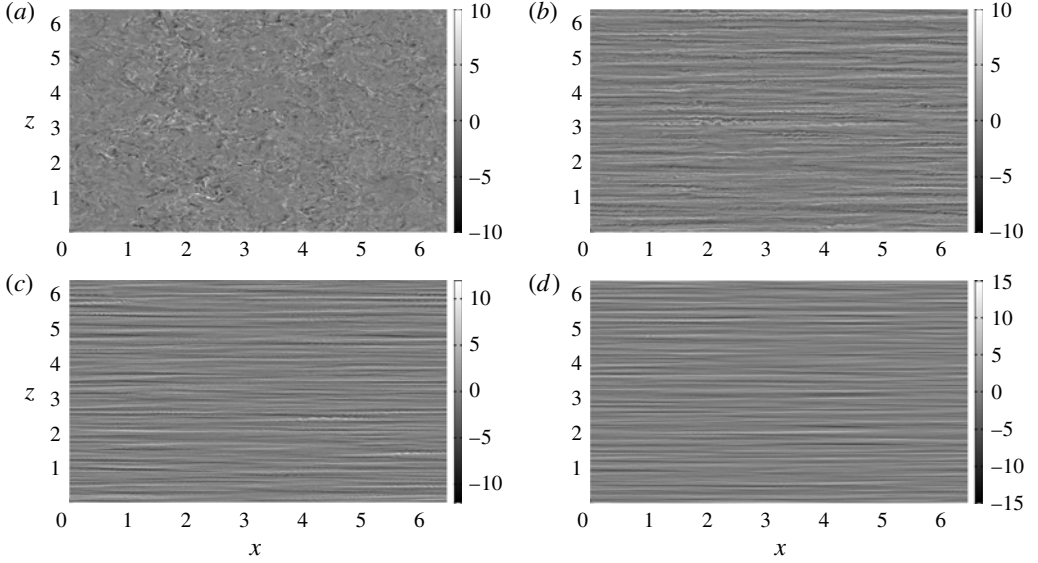


FIGURE 2. Vorticity field in the  $y$  direction  $\bar{\omega}_y$  on the  $x$ - $z$  plane at  $y=0.25$  and  $t=450$  for the high-resolution case with (a)  $N=0$ , (b)  $N=2$ , (c)  $N=4$ , and (d)  $N=6$ .

direction. KH instabilities are clearly visible in figure 2(b), but less so in figure 2(c,d); indeed, the increased stratification in figure 2(c,d) leads to decreasing numbers of KH billows. For stratified simulations, increased stratification inhibits KH instabilities as the layer thickness decreases towards the dissipation scale (e.g. Hebert & de Bruyn Kops 2006a; Brethouwer *et al.* 2007; Bartello & Tobias 2013; Khani & Waite 2014).

There is an apparent connection between cases with no inertial subrange in figure 1 and cases with no KH instabilities in figure 2. This connection has been noted previously in DNS (Waite 2014) and other kinds of LES (Khani & Waite 2014). To further investigate this connection with the dynamic Smagorinsky model, we consider the local Richardson number. The Richardson number shows the ratio of total (ambient plus perturbation) buoyancy frequency over the square of the vertical shears of horizontal motions, written as

$$Ri = \frac{N^2 - \frac{g}{\rho_0} \frac{\partial \bar{\rho}}{\partial z}}{\left(\frac{\partial \bar{u}}{\partial z}\right)^2 + \left(\frac{\partial \bar{v}}{\partial z}\right)^2}, \quad (4.1)$$

where  $N^2 = -(g/\rho_0)\partial\Phi_0/\partial z$  is constant,  $\rho_0$ ,  $g$  and  $\Phi_0$  are the reference density, gravity and ambient density fields, respectively. Small Richardson numbers including negative values correspond to overturning and KH instabilities. Figure 3 shows the time-averaged histograms of the local Richardson number  $Ri$  for the high- and low-resolution LES with different buoyancy frequencies. Histograms show long tails for positive  $Ri$  and also rapid drops for negative  $Ri$ . It appears that by increasing the resolution and decreasing buoyancy frequency, the peaks in the  $Ri$  histograms move towards negative values, i.e. more of the domain is subject to KH and gravitational instabilities. In addition, the high-resolution case with  $N=2$  (the solid black line)



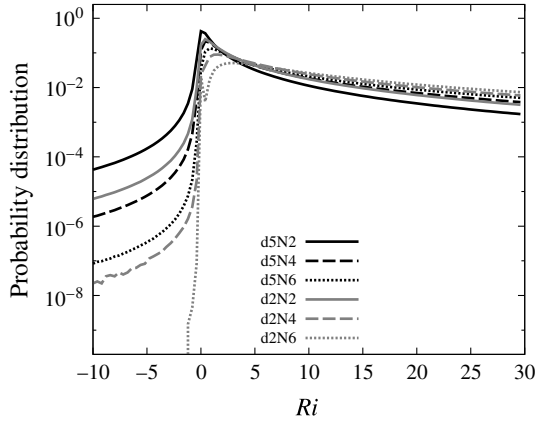


FIGURE 3. The averaged histograms of local Richardson number  $Ri$ . Only the segment  $-10 \leq Ri \leq 30$  is shown. Histograms are normalized by bin size to give probability distributions, and are computed with 100 bins over the given segment ( $\Delta Ri = 0.4$ ).

shows the largest number of points with negative values and the smallest number of points with positive  $Ri$ . However, the low-resolution case with  $N=6$  (the grey dotted line) shows the largest values for positive  $Ri$  and just a few points with negative  $Ri$ . Consistent with the Smagorinsky and Kraichnan LES of Khani & Waite (2014), increased resolution at fixed buoyancy frequency, or decreased stratification at fixed resolution, increases the number of points with small and negative  $Ri$ . As a result, figure 3 demonstrates that increased resolution or decreased stratification are more likely to lead to generation of KH instabilities and overturning (consistent with the results of figure 2).

Considering these findings, we can now attempt to find a threshold for  $\Delta/L_b$  which guarantees that the breakdown of the layers into KH instabilities is captured (see Khani & Waite 2014 for similar thresholds in the Kraichnan and Smagorinsky LES). According to table 1 and figure 3, if we consider cases with  $\Delta < 0.24L_b$  (i.e.  $k_c > 2.1k_b$ ), the averaged  $Ri$  histograms for  $-10 \leq Ri \leq 30$  are above  $10^{-6}$ . On the other hand, if  $\Delta > 0.24L_b$ , only very few points with small and negative  $Ri$  appear. These findings suggest that the maximum grid spacing for the dynamic Smagorinsky model should be in the range  $0.18L_b < \Delta < 0.24L_b$ . In contrast, the threshold for the regular Smagorinsky LES is in the range  $0.12L_b \leq \Delta < 0.17L_b$  (see Khani & Waite 2014). While the dynamic Smagorinsky values are larger – i.e. lower resolution is able to capture the basic dynamics of stratified turbulence – these ranges nearly overlap, suggesting that the criteria for the regular and dynamic Smagorinsky are close. In the next section we compare these two SGS models and show that the dynamic Smagorinsky model is clearly better.

#### 4.2. The dynamic Smagorinsky model versus the Smagorinsky model

To get a better understanding of how well the dynamic Smagorinsky model performs in stratified turbulence simulations, we compare with results obtained with the regular Smagorinsky model at the same stratifications and resolutions (results from Khani & Waite 2014; these simulations are labelled ‘S’ in table 1). Figure 4 shows the horizontal wavenumber energy spectra at different resolutions and fixed buoyancy frequency for both SGS models. Clearly, the dynamic version

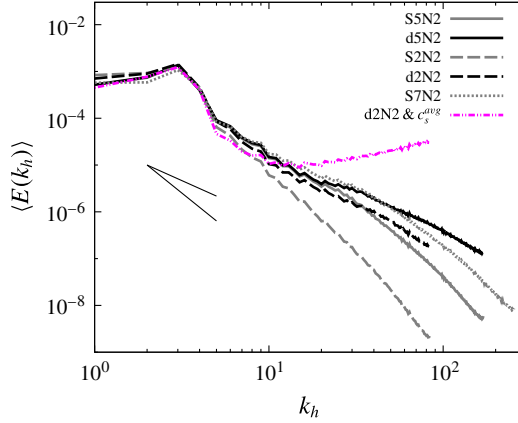


FIGURE 4. (Colour online) The averaged horizontal wavenumber energy spectra for the dynamic Smagorinsky cases compared with the Smagorinsky cases at different resolutions with fixed buoyancy frequency  $N = 2$ . The dot-dot-dashed (magenta online) curve shows the averaged horizontal wavenumber energy spectra over  $300 \leq t \leq 356$  for dynamic Smagorinsky simulation, in which  $c_s$  values are space-averaged instead of clipping negative values. The solid black line segments show  $-5/3$  and  $-3$  slopes.

is less dissipative than the regular Smagorinsky model with the same resolution at large horizontal wavenumbers: at the same resolutions, the dynamic and regular Smagorinsky results agree fairly well at large scales but diverge at small scales, where the regular Smagorinsky results exhibit a much broader and steeper dissipation range. This is particularly pronounced at the lower resolution ( $n = 256$ ), where the regular Smagorinsky spectrum is much steeper than the dynamic model over most wavenumbers. It is interesting that the dynamic case with  $n = 512$  is very similar to the Smagorinsky LES with  $n = 768$ ; nevertheless, the Smagorinsky model is still more dissipative for  $k_h \gtrsim 60$ . In addition, the low-resolution dynamic Smagorinsky case with  $N = 2$  is very close to the regular Smagorinsky LES with  $n = 512$  and  $N = 2$ . This trend implies that the dynamic Smagorinsky LES looks like the regular Smagorinsky case with twice the resolution. As a result, the low-resolution dynamic Smagorinsky model yields similar results to the high-resolution Smagorinsky model at fixed buoyancy frequency. It is worth mentioning that at the same resolution, both the dynamic and regular Smagorinsky models are almost identical at large scales: for the case with  $n = 256$ , they are very similar for  $k_h < 6$ , and for the case with  $n = 512$ , they are almost identical up to  $k_h = 20$ .

Figure 5 shows the horizontal and vertical wavenumber spectra of SGS energy transfer (i.e. eddy dissipation spectra) for stratified and unstratified dynamics models at  $t = 450$ . The eddy dissipation spectrum shows the rate of kinetic energy dissipation at resolved wavenumbers due to the eddy viscosity mechanism. In other words, the eddy dissipation spectra indicate the energy transfer from resolved scales towards the (unresolved) SGS motions. In order to compute the spectral SGS energy transfer, we multiply the Fourier coefficient of the eddy viscosity term in (2.1) with twice the complex conjugate of the Fourier coefficient of the filtered velocity field, and take the real part. For comparison, the eddy dissipation spectra of the Smagorinsky LES for the case with  $n = 768$  and  $N = 2$  (see Khani & Waite 2014) are also shown. For unstratified cases, the maximum eddy dissipation occurs at small scales for both

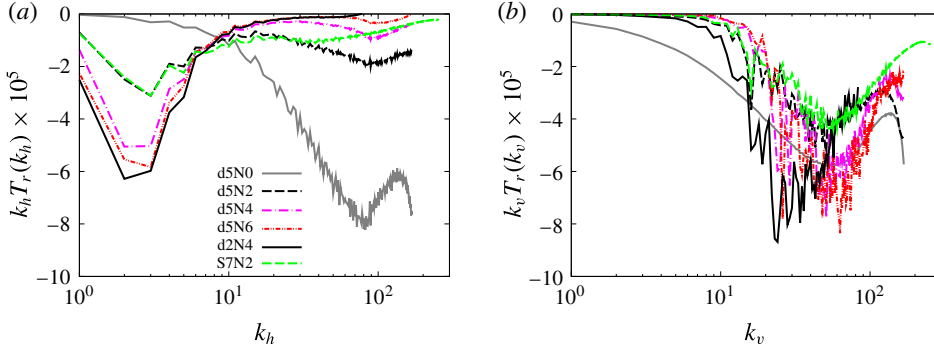


FIGURE 5. (Colour online) The horizontal (a) and vertical (b) wavenumber spectra of SGS energy transfer for the low-resolution and high-resolution cases at  $t = 450$ . The spectra are multiplied by wavenumber in order to preserve area on the log-linear axes. For comparison, the results of high-resolution Smagorinsky LES are also shown (Khani & Waite 2014).

the horizontal and vertical wavenumber spectra. These trends are consistent with the isotropic dissipation picture, in which the kinetic energy is damped mainly at small scales. For stratified cases, however, the eddy dissipation spectra are very different in the horizontal and vertical directions. According to figure 5(a), the horizontal wavenumber spectra of SGS energy transfer peak at large scales.

Unlike the horizontal wavenumber spectra, the vertical wavenumber spectra of eddy dissipation for the stratified cases peak at small scales (figure 5b). Increased stratification or decreased resolution increase the vertical SGS energy transfer spectra by contracting the thickness of vertical layers towards the dissipation scale or increasing dissipation scales, respectively (as seen in previous hyperviscosity simulations and DNS: Waite & Bartello 2004; Hebert & de Bruyn Kops 2006b; Brethouwer *et al.* 2007; Almkie & de Bruyn Kops 2012; Bartello & Tobias 2013). As a result, for an isotropic eddy viscosity model, increased stratification leads to anisotropic eddy dissipation spectra due to strongly anisotropic resolved scales (see also Khani & Waite 2014 for the Kraichnan and Smagorinsky models). It is worth mentioning that the horizontal and vertical eddy dissipation spectra of the dynamic Smagorinsky LES with  $n = 512$  and  $N = 2$  are very similar to their counterparts in the Smagorinsky LES with higher resolution  $n = 768$  and the corresponding buoyancy frequency, especially when  $k_{h,v} \lesssim 60$  (figure 5). As a result, the low-resolution dynamic Smagorinsky model generates almost the same eddy dissipation as the high-resolution Smagorinsky model in LES of stratified turbulence. The results in figure 5 are consistent with the results of Khani & Waite (2013), who measured effective eddy viscosity and dissipation spectra in DNS of stratified turbulence. In particular, the maximum energy transfer from resolved scales towards SGS motions happens at large horizontal and small vertical scales (in line with the anisotropic dissipation in DNS of Khani & Waite 2013).

Figure 6 shows the horizontal and vertical energy transfer spectra for the stratified case with  $n = 512$  and  $N = 2$  at  $t = 450$ . The total kinetic energy transfer spectrum, which is shown by a solid black line in figure 6, is a summation of the resolved energy transfer spectrum  $T^*$  plus the eddy dissipation spectrum  $T_r$ . In the horizontal direction (figure 6a), there is a large negative peak, which balances the energy input at large horizontal scales by the forcing term. In the intermediate range  $7 \lesssim k_h \lesssim 23$ ,

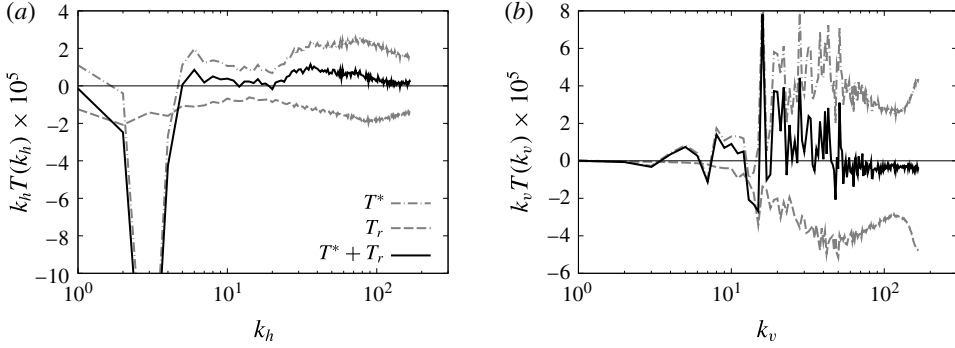


FIGURE 6. The horizontal (a) and vertical (b) energy transfer spectra for the high-resolution stratified case with  $N = 2$  at  $t = 450$ . The spectra are multiplied by wavenumbers in order to preserve area on the log-linear axes. The thin black line shows the zero value.

the energy transfer is approximately zero, which shows the stratified turbulence inertial subrange (Lindborg 2006). At smaller resolved scales a positive bump around  $k_h = 30$  is seen, which indicates the injection of energy near the buoyancy scale  $k_b$ , consistent with the development of KH instabilities and billows (in line with DNS and hyperviscosity simulations of Waite 2011, 2014). Figure 6(b) shows the vertical energy transfer spectrum, in which a peak around the buoyancy scale  $k_b = 30$  is observed, as in the horizontal. Overall, the transfer spectra seem broadly consistent with DNS results, which show transfer of energy out of the forcing scales and into the buoyancy scale (Almalkie & de Bruyn Kops 2012; Waite 2014).

In conclusion, this section shows that the maximum criterion on  $\Delta/L_b$  for resolving KH instabilities in the dynamic Smagorinsky model is definitely larger than that for the regular Smagorinsky model. Even with half the resolution, the dynamic Smagorinsky model gives similar results to the regular Smagorinsky model (figure 4).

#### 4.3. The dynamic Smagorinsky coefficient $c_s$

In this section, a detailed analysis on the dynamic Smagorinsky coefficient  $c_s$  is presented. Figure 7(a) shows a snapshot for the  $\bar{c}_s^+$  field in the  $x$ - $z$  plane along with a few contours of the characteristic rate of strain  $\bar{S} = (2\bar{s}_{ij}\bar{s}_{ij})^{1/2}$  at  $y = 0.25$  and  $t = 450$  for the high-resolution case with  $N = 4$ . The vertical axis is zoomed in to show around 13 vertical layers of length  $2\pi/k_b$ . It is interesting that the presence of stratification leads to vertical layers in the  $\bar{c}_s^+$  field. Black contours of  $\bar{S}$  indicate low values  $\bar{S} = 1$  and 1.5. The magenta contours on the other hand show higher values  $\bar{S} = 6$  and 8. Interestingly, regions with high straining in figure 7 are generally associated with zero or very small values of  $\bar{c}_s^+$ , while regions with low straining correspond to larger values of  $\bar{c}_s^+$ . This behaviour, which has been investigated elsewhere in unstratified turbulence (e.g. Kleissl *et al.* 2006; Wan & Porté-Agel 2011), suggests that the dynamic Smagorinsky coefficient  $c_s$  decreases as shear increases in order to preserve the small-scale instabilities between vertical layers.

Figure 7(b) shows the probability distributions of  $\bar{c}_s^+$  on gridpoints restricted to different ranges of  $\bar{S}$  at  $t = 450$ . It is clear that increased straining is associated with smaller values of  $\bar{c}_s^+$ . With weak straining, distributions of the dynamic Smagorinsky

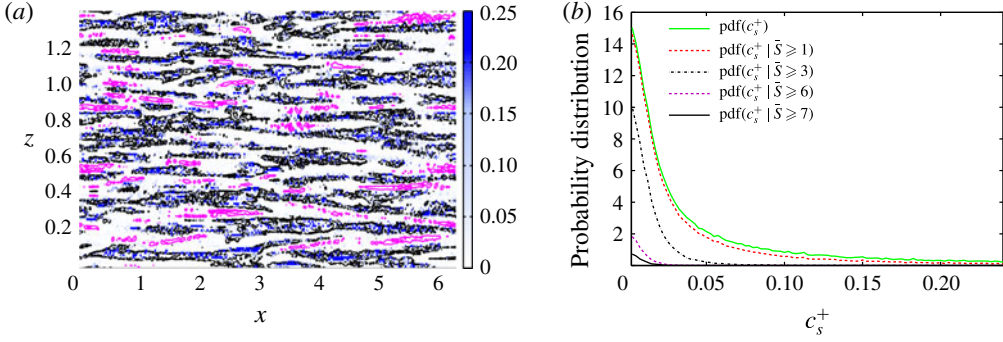


FIGURE 7. (a) The Smagorinsky coefficient field  $c_s^+$  on the  $x$ - $z$  plane and (b) the conditional distribution of  $c_s^+$ , for the high-resolution case with  $N = 4$  at  $y = 0.25$  and  $t = 450$ . In (a), contours of constant  $\bar{S} = (2\bar{s}_{ij}\bar{s}_{ij})^{1/2}$  are overlaid on the  $c_s^+$  field where the vertical axis is zoomed in to  $80/k_b$  that includes around  $13L_b$ . The magenta contours present high values  $\bar{S} = 6$  and  $8$  and the black contours show those of low values  $\bar{S} = 1$  and  $1.5$ .

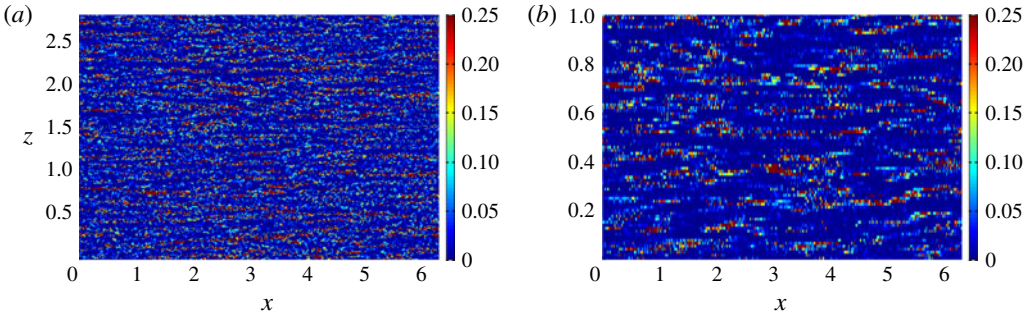


FIGURE 8. The Smagorinsky coefficient field  $c_s^+$  on the  $x$ - $z$  plane for the high-resolution cases with (a)  $N = 2$  and (b)  $N = 6$  at  $y = 0.25$  and  $t = 450$ , where the vertical axes are zoomed in to  $80/k_b$  to include around  $13L_b$ .

coefficient  $c_s^+$  show both small and large  $c_s^+$  values. However, for large straining, these distributions show small  $c_s^+$  values. For example, for  $\bar{S} \geq 7$ , the solid black line in figure 7(b),  $c_s^+$  values are smaller than  $(0.14)^2$ . This trend is consistent with the suggestion of Deardorff (1971) that the presence of shear implies small  $c_s^+$  values. In addition, looking at  $c_s^+$  fields at different stratifications shows that increased stratification at fixed resolution leads to vertical layers that are not fully resolved, consistent with failure of the criterion on  $\Delta/L_b$  (see figure 8).

The time-averaged histograms of  $c_s$  at different resolutions and buoyancy frequencies are shown in figure 9(a). Increased stratification leads to decreasing standard deviations, i.e.  $c_s$  values tend to get smaller (consistent with figure 7). Interestingly, the distribution of negative  $c_s$  (before clipping is applied) is very similar to that of positive  $c_s$ ; similar trends are seen in Kang, Chester & Meneveau (2003), Meneveau, Lund & Cabot (1996) for probability distributions of local SGS dissipation, and probability distributions of  $c_s$ , respectively. As discussed in §2, negative  $c_s$  values occur when the resolved stress  $L_{ij}$  and Germano's rate of strain  $M_{ij}$  are not coaxial, implying backscatter. Since we remove negative  $c_s$  values at every time step to prevent numerical instabilities, the clipping procedure removes some local information about

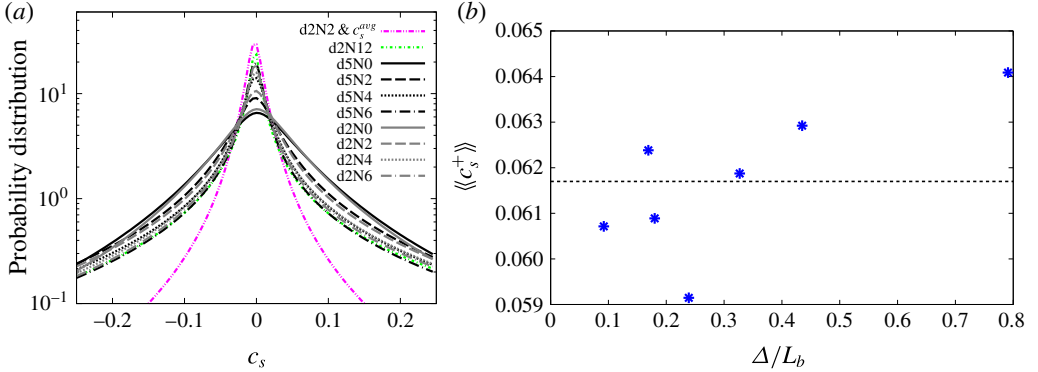


FIGURE 9. (Colour online) (a) The averaged histograms of Smagorinsky coefficient  $c_s$  and (b) the time- and space-averaged  $c_s$  versus  $\Delta/L_b$ . In (a), only the segment  $-0.25 \leq c_s \leq 0.25$  is shown. Histograms are normalized by bin size to give probability distributions, and are computed with 100 bins over the given segment ( $\Delta c_s = 0.005$ ). The dashed line in (b) shows the average of all  $\langle\langle c_s^+ \rangle\rangle$  values.

the modelled SGS stress tensor and the dynamics of upscale energy transfer, which are not appropriately represented with an eddy viscosity model. In addition, this plot shows that the volume-averaged  $c_s$  values are very close to zero. Figure 9(b) shows the time- and space-averaged  $c_s$  values versus  $\Delta/L_b$ . It is interesting that the time- and space-averaged  $c_s^+$  for stratified cases is around  $\langle\langle c_s^+ \rangle\rangle = (0.25)^2$  (the dashed black line in figure 9b), which is only slightly larger than that of the non-stratified case with  $\langle\langle c_s^+ \rangle\rangle = (0.23)^2$ , and also larger than the regular Smagorinsky constant  $c_s = (0.17)^2$ . This behaviour might suggest that the presence of stratification decreases straining in turbulent flows and so the dynamic Smagorinsky model implies larger  $c_s$  in stratified turbulence, corresponding to changes in dynamics of flows as stratification appears (see also figure 7). Interestingly, the probability distributions of  $c_s$  in stratified turbulence are symmetric, which is not the case for  $c_s^+$  (in line with the isotropic turbulence: Meneveau *et al.* 1996; Kang *et al.* 2003).

It is interesting that the self-similar assumption of the dynamic Smagorinsky model works reasonably well in stratified turbulence when  $\Delta \sim L_b$ . A possible explanation for this behaviour could be the development and breakdown of KH billows, which are much more isotropic than the larger-scale layerwise structures. In addition, DNS results of stratified turbulence show that the effective spectral eddy viscosity has self-similar structure when the test cutoff  $k_c$  is around  $k_o$  or smaller (Khani & Waite 2013).

An alternative to clipping negative  $c_s$  values is to use spatial averaging of  $c_s$  over the homogeneous directions, as is often done in boundary layer simulations (e.g. Meneveau & Katz 2000; Pope 2000; Porté-Agel *et al.* 2000). Here, since we are considering homogeneous stratified turbulence, this approach amounts to taking a spatial average of  $c_s$  before evaluating the eddy viscosity (1.2). We have evaluated this approach by performing an additional simulation with  $n = 256$  and  $N = 2$ , in which  $c_s$  values are averaged over the entire domain instead of clipping negative  $c_s$  values. Because of the lack of sufficient dissipation at small scales, the simulation in this case blew up at around  $t = 356$  (see the dot-dot-dashed (magenta online) curve in figure 4, which clearly shows an artificial pile-up of energy at large wavenumbers). The averaged probability distribution of  $c_s$  for this simulation over  $300 \leq t \leq 356$  is



also shown in figure 9(a), in which shorter tails and a larger peak are seen. For this case, the standard deviation is half of that seen when clipping negative  $c_s$  values.

## 5. Conclusions

A systematic analysis of the performance of the dynamic Smagorinsky model in stratified turbulence has been studied in this paper. Our results show that if  $\Delta/L_b < 0.24$ , several key features of stratified turbulence – including an energy spectrum with an approximately  $-5/3$  slope in the horizontal, KH instabilities and overturning – are captured in LES of stratified turbulence using the dynamic Smagorinsky model. Other SGS models have different criteria for  $\Delta/L_b$ . For example, Khani & Waite (2014) argued that for the Smagorinsky and Kraichnan model  $\Delta/L_b$  should be less than 0.17 and 0.47, respectively, to ensure these features of stratified turbulence are captured. As a result, the resolution limit on  $\Delta$  in the dynamic Smagorinsky model is 40 % larger than that of the Smagorinsky model (see figure 4). In addition, this study shows that spectra of SGS energy transfers in stratified simulations are significantly different in the horizontal and vertical directions, in line with other LES models (Khani & Waite 2014). The maximum eddy dissipation spectra in stratified cases occur at large horizontal and small vertical scales. This behaviour suggests that spectra of SGS energy transfer are anisotropic in stratified turbulence. Overall, the eddy dissipation from the dynamic Smagorinsky LES with  $N = 2$  and  $n = 512$  looks similar to the regular Smagorinsky LES at the same buoyancy frequency with  $n = 768$ .

The dynamic eddy viscosity coefficient  $c_s^+$  shows a layerwise field in the presence of stratification, consistent with the familiar layered structure of stratified turbulence. This finding implies that the dynamic Smagorinsky model can make a connection between the SGS model and stratification and that this connection yields an improvement on the performance of traditional SGS models like the Smagorinsky model in stratified turbulence. It is interesting that large values of  $c_s^+$  correspond to regions with weak straining. Consistently, in regions with large shears (strong vertical straining), the values of  $c_s^+$  are very small. This behaviour explains why the dynamic Smagorinsky model is overall less dissipative than the regular Smagorinsky model. Also, the overall relationship between  $c_s^+$  and shear  $\tilde{S}$  is similar to what has been discussed in the literature on unstratified turbulence. In addition, local distributions of negative  $c_s$  are very close to positive  $c_s$ , implying that local negative dissipation (i.e. backscatter) might be important in the dynamic Smagorinsky model. Nevertheless, negative values of  $c_s$  are set to zero in this study to stabilize the simulations and ensure sufficient small-scale dissipation. Ultimately, the appropriate inclusion of backscatter in LES requires a different approach from the eddy viscosity model (e.g. Carati, Ghosal & Moin 1995; Ghosal *et al.* 1995).

At the same resolution, the dynamic Smagorinsky model is much more expensive than the regular Smagorinsky model; wall clock run times are approximately twice as long. This increase in run time is due to the significant increase in the number of fast Fourier transforms (FFTs) required to compute  $c_s$ . On the other hand, the dynamic Smagorinsky model can be run at lower resolution and still give similar results to the regular Smagorinsky model. Decreasing the resolution by 33 % (i.e. from  $n = 768$  to  $n = 512$ ) cuts the run time by a factor of 5. Halving the resolution (i.e. from  $n = 512$  to  $n = 256$ ) cuts the run time by a factor of 16, and still gives dynamic Smagorinsky results similar to regular Smagorinsky. Despite the increased number of FFTs, the ability to run at lower resolution implies that the dynamic Smagorinsky approach is clearly more efficient than the regular Smagorinsky method.



Overall, we have now looked at three SGS models: Kraichnan, Smagorinsky and dynamic Smagorinsky. Despite the fact that these models were designed for unstratified turbulence, they work well for stratified turbulence when  $L_b$  is resolved sufficiently. The interpretation of ‘sufficiently’ depends on the model. The Kraichnan model is clearly the best – i.e. it requires the lowest resolution to give the same results – but it only works for triply periodic spectral models. Otherwise, the dynamic Smagorinsky model is better than regular Smagorinsky. However, none of these models work well when  $L_b$  is not resolved. Eliminating this barrier, i.e. running LES without fully resolving  $L_b$ , would require a major rework. For future work, performing LES of stratified turbulence with anisotropic eddy viscosity terms could be considered. In addition, considering local backscatter in the dynamic SGS models is another potential avenue in studying of stratified turbulence. Ultimately, we need to perform a high-resolution DNS of stratified turbulence to study the dynamics of energy transfer around the buoyancy scale  $L_b$  and maybe the Ozmidov scale  $L_o$ .

### Acknowledgements

Computations were performed on the GPC supercomputer at the SciNet HPC Consortium. SciNet is funded by the Canada Foundation for Innovation under the auspices of Compute Canada, the Government of Ontario, Ontario Research Fund – Research Excellence, and the University of Toronto. Also, this work was made possible by the facilities of the Shared Hierarchical Academic Research Computing Network (SHARCNET: [www.sharcnet.ca](http://www.sharcnet.ca)) and Compute/Calcul Canada. Financial support from the Natural Sciences and Engineering Research Council of Canada is gratefully acknowledged.

### REFERENCES

- ALMALKIE, S. & DE BRUYN KOPS, S. M. 2012 Kinetic energy dynamics in forced, homogeneous, and axisymmetric stably stratified turbulence. *J. Turbul.* **13**, 1–32.
- BARTELLO, P. & TOBIAS, S. M. 2013 Sensitivity of stratified turbulence to the buoyancy Reynolds number. *J. Fluid Mech.* **725**, 1–22.
- BRETHOUWER, G., BILLANT, P., LINDBORG, E. & CHOMAZ, J.-M. 2007 Scaling analysis and simulation of strongly stratified turbulent flows. *J. Fluid Mech.* **585**, 343–368.
- CARATI, D., GHOSAL, S. & MOIN, P. 1995 On the representation of backscatter in dynamic localization models. *Phys. Fluids* **7** (3), 606–616.
- DEARDORFF, J. W. 1971 On the magnitude of the subgrid scale eddy coefficient. *J. Comput. Phys.* **7**, 120–133.
- DOMARADZKI, J. A., LIU, W. & BRACHET, M. E. 1993 An analysis of subgrid-scale interactions in numerically simulated isotropic turbulence. *Phys. Fluids A* **5** (7), 1747–1759.
- GERMANO, M. 1992 Turbulence: the filtering approach. *J. Fluid Mech.* **238**, 325–336.
- GERMANO, M., PIOMELLI, U., MOIN, P. & CABOT, W. H. 1991 A dynamic subgrid-scale eddy viscosity model. *Phys. Fluids A* **3** (7), 1760–1765.
- GHOSAL, S., LUND, T. S., MOIN, P. & AKSELVOLL, K. 1995 A dynamic localization model for large-eddy simulation of turbulent flows. *J. Fluid Mech.* **286**, 229–255.
- HEBERT, D. A. & DE BRUYN KOPS, S. M. 2006a Relationship between vertical shear rate and kinetic energy dissipation rate in stably stratified flows. *Geophys. Res. Lett.* **33**, L06602.
- HEBERT, D. A. & DE BRUYN KOPS, S. M. 2006b Predicting turbulence in flows with strong stable stratification. *Phys. Fluids* **18**, 066602.
- HERRING, J. R. & MÉTAIS, O. 1989 Numerical experiments in forced stably stratified turbulence. *J. Fluid Mech.* **202**, 97–115.

- JIMÉNEZ, J. R. & MOSER, R. D. 2000 Large-eddy simulations: where are we and what can we expect? *AIAA J.* **38** (4), 605–612.
- KANG, H. S., CHESTER, S. & MENEVEAU, C. 2003 Decaying turbulence in an active-grid-generated flow and comparisons with large-eddy simulation. *J. Fluid Mech.* **480**, 129–160.
- KHANI, S. & WAITE, M. L. 2013 Effective eddy viscosity in stratified turbulence. *J. Turbul.* **14** (7), 49–70.
- KHANI, S. & WAITE, M. L. 2014 Buoyancy scale effects in large-eddy simulations of stratified turbulence. *J. Fluid Mech.* **754**, 75–97.
- KLEISSL, J., KUMAR, V., MENEVEAU, C. & PARLANGE, M. B. 2006 Numerical study of dynamic Smagorinsky model in large-eddy simulation of the atmospheric boundary layer: validation in stable and unstable conditions. *Water Resour. Res.* **42**, W06D10.
- KRAICHNAN, R. H. 1976 Eddy viscosity in two and three dimensions. *J. Atmos. Sci.* **33**, 1521–1536.
- LAVAL, J.-P., MCWILLIAMS, J. C. & DUBRULLE, B. 2003 Forced stratified turbulence: successive transition with Reynolds number. *Phys. Rev. E* **68**, 036308.
- LESIEUR, M. & MÉTAIS, O. 1996 New trends in large-eddy simulations of turbulence. *Annu. Rev. Fluid Mech.* **28**, 45–82.
- LILLY, D. K. 1967 The representation of small-scale turbulence in numerical simulation experiments. In *NCAR Manuscript 281*, National Center for Atmospheric Research, Boulder, CO, USA, pp. 99–164.
- LILLY, D. K. 1992 A proposed modification of the Germano subgrid-scale closure method. *Phys. Fluids A* **4** (3), 633–635.
- LINDBORG, E. 2006 The energy cascade in strongly stratified fluid. *J. Fluid Mech.* **550**, 207–242.
- LU, H. & PORTÉ-AGEL, F. 2014 On the development of a dynamic nonlinear closure for large-eddy simulation of the atmospheric boundary layer. *Boundary-Layer Meteorol.* **151** (3), 429–451.
- MENEVEAU, C. 2012 Germano identity-based subgrid-scale modeling: a brief survey of variations on a fertile theme. *Phys. Fluids* **24**, 121301.
- MENEVEAU, C. & KATZ, J. 2000 Scale-invariance and turbulence models for large-eddy simulation. *Annu. Rev. Fluid Mech.* **32**, 1–32.
- MENEVEAU, C., LUND, T. S. & CABOT, W. H. R. 1996 A Lagrangian dynamic subgrid-scale model of turbulence. *J. Fluid Mech.* **319**, 353–385.
- ORSZAG, S. A. 1971 On the elimination of aliasing in finite-difference schemes by filtering high-wavenumber components. *J. Atmos. Sci.* **28**, 1074–1074.
- PAOLI, R., THOURON, O., ESCOBAR, J., PICOT, J. & CARIOLLE, D. 2013 High-resolution large-eddy simulations of sub-kilometer-scale turbulence in the upper troposphere lower stratosphere. *Atmos. Chem. Phys. Discuss.* **13**, 31891–31932.
- PIOMELLI, U. 1999 Large-eddy simulation: achievements and challenges. *Prog. Aerosp. Sci.* **35**, 335–362.
- POPE, S. B. 2000 *Turbulent Flows*. Cambridge University Press.
- POPE, S. B. 2004 Ten questions concerning the large-eddy simulation of turbulent flows. *New J. Phys.* **6** (35), 1–24.
- PORTÉ-AGEL, F., MENEVEAU, C. & PARLANGE, M. B. 2000 A scale-dependent dynamic model for large-eddy simulation: application to a neutral atmospheric boundary layer. *J. Fluid Mech.* **415**, 261–284.
- REMMLER, S. & HICKEL, S. 2012 Direct and large eddy simulation of stratified turbulence. *Intl J. Heat Fluid Flow* **35**, 13–24.
- RILEY, J. J. & DE BRUYN KOPS, S. M. 2003 Dynamics of turbulence strongly influenced by buoyancy. *Phys. Fluids* **15**, 2047–2059.
- SIEGEL, D. A. & DOMARADZKI, J. A. 1994 Large-eddy simulation of decaying stably stratified turbulence. *J. Phys. Oceanogr.* **24**, 2353–2386.
- SMAGORINSKY, J. 1963 General circulation experiments with the primitive equations. Part I. The basic experiment. *Mon. Weath. Rev.* **91** (3), 99–164.
- SMITH, C. M. & PORTÉ-AGEL, F. 2014 An intercomparison of subgrid models for large-eddy simulation of katabatic flows. *Q. J. R. Meteorol. Soc.* **140** (681), 1294–1303.
- WAITE, M. L. 2011 Stratified turbulence at the buoyancy scale. *Phys. Fluids A* **23**, 066602.

- WAITE, M. L. 2014 Direct numerical simulations of laboratory-scale stratified turbulence. In *Modeling Atmospheric and Oceanic Flows: Insights from Laboratory Experiments and Numerical Simulations* (ed. T. von Larcher & P. Williams), American Geophysical Union, pp. 159–175. Wiley & Sons.
- WAITE, M. L. & BARTELLO, P. 2004 Stratified turbulence dominated by vortical motion. *J. Fluid Mech.* **517**, 281–303.
- WAN, F. & PORTÉ-AGEL, F. 2011 Large-eddy simulation of stably-stratified flow over a steep hill. *Boundary-Layer Meteorol.* **138**, 367–384.